Tracking the local seismicity level in the active influence zone of the southern Uzbekistan reservoirs

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Abstract. Assessment of local seismicity caused by human engineering activity is a necessity after several earthquakes have been recorded in world practice, which destroyed a hydraulic structure, a dam and accompanied by human casualties.

The Tupalang, Gissarak and Pachkamar reservoirs, operated in southern Uzbekistan, are located in the mountain system of the southern part of the Tien Shan seismogenic fault.

According to the general seismic zoning, the zones where these reservoirs are located are classified as 8-9 magnitude shakes during strong earthquakes. Localisation of local earthquakes associated with the operation of these reservoirs is necessary for taking operational measures to prevent possible emergencies associated with an increase in the level of technogenic seismicity.

Methods for assessing the local seismic regime were used to track the level of local seismicity. The catalogs of local earthquakes occurring in the zone of active influence of reservoirs have been compiled.

The seismic activity levels and changes in the angle of inclination of the recurrence of earthquakes were determined for each object separately. The dependence of the variation in the level of local seismicity on the mode of operation of reservoirs has been revealed.

Several evidence of the negative impact of large hydraulic reservoirs on the deformation and seismic state of the zones of influence of reservoirs in southern Uzbekistan have been identified.

1 Introduction

In the complex of studies that contribute to ensuring the geodynamic safe operation of reservoirs in Uzbekistan, digital seismic monitoring is becoming a new direction [1]. It is based on the representative registration of seismic vibrations from small acts of brittle fracture in the side rock mass [2, 3].

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By processing seismic records, the characteristics of destruction processes are assessed -
their position in space, released seismic energy, inelastic deformation in the source area [4].
The potential benefits of quantitative processing and interpretation of seismic events in the
basin of reservoirs have been known for a long time [5, 6].

However, previous attempts to organise seismic monitoring in a planned mode in the
reservoir zones of Uzbekistan were not successful, mainly due to technical limitations [7, 8,
9].

In the present, a number of new, interesting results have been obtained [1, 10, 11, 12].
Studies of seismic activity, changes in the maximum of energy classes in the areas of
reservoirs of southern Uzbekistan, calculated from representative seismological data, have
shown that the possibility of the existence of signs of technogenic influence on the
geodynamics of the area and the background seismicity of the region is quite comparable.

Therefore, the isolation from the general seismicity of the background and its local
technogenic part, the determination of the spatial grouped seismic events, which is
characterised by the establishment of zones of interrelated local seismic events, possibly
depending on the mode of operation of reservoirs, turned out to be quite in demand [11, 12,
13, 14].

The seismicity of the reservoir areas has been raised in several publications [2, 10, 16,
17, 18]. It has two main aspects [1, 2, 6, 19, 20].

First, it is important to know the degree of natural tectonic activity in this area [10, 21].

It is necessary to determine the large earthquakes that have arisen here earlier, even
before the start of the operation of reservoirs [17-19]. When analysing the catalog of
earthquakes in Uzbekistan, the catalog of strong earthquakes in Eurasia, the map of the
general seismic zoning of Uzbekistan, the answer to this question will be negative because
special studies of local seismicity with instrumental means in the reservoir areas were

carried out selectively [7-], [21].

However, it should be noted that several seismic stations (s/s) have been operating in
the areas of intensive exploitation of reservoirs in Uzbekistan for more than 40 years (s / s
Charvak - and. Charvak, s / s Kampiravat - and. Andijan, s/s Shabad - vodka. Gissarak and
others) [8,9]. But their materials were not involved in the compilation of local catalogs of
earthquakes, including weak earthquakes.

Over the past decades, these stations have been working to control the stability of the
dams of the respective reservoirs [7, 8, 9, 10]. Their materials were practically not used in
assessing local seismicity and the state of local geodynamics [10].

The first information that off-background earthquakes occur in the reservoirs of
Uzbekistan was obtained in 1974 [7, 8]. A detailed analysis of records of seismic events by
experts in assessing the stability of dams made it possible to say that several microseismic
events in the Charvak and Andijan reservoirs are, in fact, micro tectonic natural earthquakes
[8, 12].

The fact of registration of such events suggests that local earthquakes can occur in the
zone of deformation effect of reservoirs [1, 10, 12]. They depend mainly on changes in the
volume of reservoirs and the geological location of reservoirs.

The issue of local seismicity is important for a detailed specification of the seismic
hazard of areas where reservoirs with high and mainly Earth dams are located in the
mountainous area of southern Uzbekistan [8, 10, 22].

The second aspect of the problem of seismicity in the area of reservoirs is the possible
induced seismicity in this area, which is a consequence of the intensive operation of the
reservoir with a rather sharp change in their volumes in a short time [10, 16, 17, 22].

Several researchers believe that the operation of reservoirs will provoke an increase in
seismic activity even at a great distance from it [2, 3, 8, 18, 19]. There is an opinion that
rejects this possibility [5, 16, 18, 20]. However, there is still no clear data to answer this question.

2 Methods

The analysis of known cases of excited earthquakes caused by the operation of reservoirs, described in [1, 10, 11, 12, 22], made it possible to establish several empirical generalisations, taking into account the diversity of conditions of the geological structure of various reservoirs of southern Uzbekistan. Most epicenters are grouped into zones controlled by the structural features of a particular area:

- Distribution of faults, fracture zones, the strike of structures, etc. [11, 22]. The existing regional network of seismic stations in the Republic of Uzbekistan allows receiving of data on earthquakes K ≥ 9, where K is the energy class.

To obtain the necessary information, in the areas where man-made objects are located, it is necessary to achieve registration of earthquakes of lower energy classes (at least from K ≥ 7).

On the other hand, it should be taken into account that the determination of the seismic activity variation A10 and the slope of the recurrence graph is based on the recurrence of earthquakes as a fundamental property of the seismic process.

Their use is associated, in particular, with the sometimes observed decrease in the slope of the recurrence graph before strong events [12, 23].

The recurrence slope is calculated from the continuous exponential distribution of earthquakes by class. In this case, the following generally known formulas are used [1, 9, 10]:

\[
y = \log_0 \frac{1}{N} \sum\frac{K_i}{N - K_0} \tag{1}
\]

root mean square error in a value

\[
S_y = \frac{y}{\sqrt{N}} \tag{2}
\]

seismic activity A10

\[
A_{10} = \frac{1000 \cdot N \cdot (1 - 10^{-7}) \cdot 10^{-(y, S, K_0)}}{S \cdot T} \tag{3}
\]

where S is the surface area of the selected area ([S] = km²), T - duration ([T] = years).

Analysis of variations is possible only with a sufficiently large number of earthquakes. at \( y = 0.4 \) and \( N = 150 \), the standard deviation \( S_y \) is 0.04, that is, 9%.

Seismic activity A10 is calculated from the number of earthquakes N and the slope of the recurrence graph and is reduced to an area of 103 km² (for the near zone of reservoirs in terms of the area of influence) and a time interval of one year.
Around the Gissarak, Pachkamar, and Tupalang reservoirs of South Uzbekistan, a network of 7 local and 3 regional seismic stations have been set up to provide the possibility of registering earthquakes with energy classes $K \geq 7$.

3 Results and Discussion

The Gissarak reservoir (Kashkadarya region) is located in the middle course of the Aksu River, which originates in the mountains on the northern slopes of the Gissar mountain systems.

The valley of the river, where the reservoir is located, is characterised by steep sides, large slopes of channels, relatively little preservation of terrace surfaces, and a weak distribution of young Quaternary deposits [22].

Historically, it is known that strong and destructive earthquakes occurred within a radius of up to 100 km in the area of the Gissarak reservoir, such as the Karatagskoe earthquake of 1907 with a magnitude of 9-10 on the MSK scale (magnitude $M = 7.3$; the epicenter is 70 km), Baysun, 1935, with a magnitude of 8-9 points (40-60 km south of the reservoir, with $M = 6.2$) and the Chuyanchinskoye 1911 with a force of 8 points (90-100 km east of the reservoir, with $M = 6.1$).

Seismic studies carried out earlier in the period 1973-1981 and in subsequent years showed that the territory of southern Uzbekistan is characterised by high seismicity, where earthquakes of magnitude 8 ($M \geq 7$) can occur [11, 22].

The base of the dam of the Gissarak reservoir is located in a single tectonic block in a seismicity zone of 8 points [21]. Therefore, the placement of seismic receivers in this area is designed in such a way as to provide a three-dimensional environment of the volume of rocks of interest in terms of assessing stability [4, 19].

The distance between the receivers of seismic waves is about 100-500 meters, while the observations cover the volume of rocks with a maximum transverse dimension of 400-1100 m.

There were eight measuring points in the reservoir area. The seismic station at the base of the dam was in continuous recording mode. With the help of 24 receivers, a database of digital records was compiled.

The main station was located in an adit 0.5 km from the dam. Separate records during the growing season were made on the left side of the reservoir bowl and the dam. Also, two measuring points worked in the localisation mode of local fluctuations.

More than 450 records were made, of which 34 were related to high-frequency vibrations from earthquakes. The synchronous recording was carried out on the dam's crest and in the zone of the beginning of water entry into the reservoir (12 km from the dam site).

Variation of non-stationary two measuring points with translation according to the established time scale and the profiles of the starboard side of the bowl was carried out with the participation in the processing of the above-mentioned stationary stations. Varying the measuring points of observation (points - starboard), analog stations worked in the presence of one horizontal component of the record.

There weren't enough of them. Therefore, the method of the non-stationary shift of measuring points was used [24]. The highest density of seismic events is observed in the southeastern part of the reservoir (southern Tien Shan). They are also partly found in the North Afghan geodynamic test site (Northern Pamir).

The lowest density is in the western part of the territory. Table 1 below shows the station detection thresholds compared to others, including the opening date of stations at various observation points.
Table 1. Comparison of useful signal detection thresholds for engineering seismometric channels (for Gindikush and local earthquakes with $M = 3.2$-$4.5$ according to the IRIS Energy Catalog)

<table>
<thead>
<tr>
<th>Observation point</th>
<th>The code</th>
<th>Station opening date</th>
<th>N-S mB</th>
<th>N-S mB</th>
<th>E-W mB</th>
<th>E-W mB</th>
<th>Z mB</th>
<th>Z mB</th>
</tr>
</thead>
<tbody>
<tr>
<td>Gissarak</td>
<td>GSR</td>
<td>26.10.2014</td>
<td>4.0</td>
<td>3.9</td>
<td>1.3</td>
<td>1.4</td>
<td>0.5</td>
<td>0.4</td>
</tr>
<tr>
<td>Pachkamar</td>
<td>PCHK</td>
<td>06.03.2015</td>
<td>3.8</td>
<td>3.7</td>
<td>4.0</td>
<td>3.9</td>
<td>0.3</td>
<td>0.3</td>
</tr>
</tbody>
</table>

For each seismometric channel of the station, detection thresholds are also shown in mV (millivolts), exceeding any of them under the action of seismic vibrations transferred the station from the standby mode to the “earthquake recording” mode.

Over the entire observation period, the seismic workstations CPC KM/V-F from/s "Shabat" with the group of seismic stations "Gissarak" and "Pachkamar" continued to stably record seismic events. The group of s / s "Tupalang" located in the near the Tupalang reservoir is also included in the processing.

All these stations jointly provided the ability to register earthquakes with $K \leq 7$ energy class occurring in southwestern Uzbekistan.

To study seismicity, the parameters of seismic sources were preliminarily determined: time in the source, coordinates of epicenters, depths of hypocenters and magnitudes, and replenishment of the local catalog of earthquakes in the near zones of the Gissarak reservoir.

Based on the distribution of earthquakes, an assessment of their parameters (seismic activity, maximum earthquake magnitude), it is possible to identify potential seismic zones by contours and assess the level of seismic risk.

The monitoring of seismicity parameters of local areas and near zones of the Gissarak reservoir includes an assessment of the level of regional seismicity according to the current scale; construction of schemes of areal distribution of local and regional earthquakes; A10 activity; the slope of the earthquake recurrence graph.

The local catalog was compiled for energy classes $K \leq 7$ (magnitude $M \geq 2$) and included more than 500 earthquakes that occurred from 2008 to mid-2016 within a radius of up to 100 km from the dam site [1].

When studying the seismicity of the Gissarak reservoir zone, weak local earthquakes with $K = 7$ were separated. The manifestation of the above shaking has increased since 2008 on records of seismic stations operating in standby mode and since 2010 on digital records.

Weak seismic shocks intensified in 2011 and 2013 and by the middle of 2013 and 2014 when the water level in the reservoir rose above the $H = 80$ m marks [11].

From the end of 2014 to August 2015, more than 30 seismic events were recorded. As for the slope of the repeatability graph $\gamma$, the discrepancy lies within the error limits. For the average long-term value of the slope of the recurrence graph for the Gissarak region, $\gamma = 0.57 \pm 0.07$ should be left.

For the last two time intervals, the value of $\gamma$ of the repeatability graph has sharply decreased. During this period, the filling of the reservoir began, and in the shortest time after that, the frequency of changes in the volume of water increased, which probably led to a change in $\gamma$ (Table 2).
Table 2. Value of the parameters of the recurrence graph of the Gissarak region

<table>
<thead>
<tr>
<th>№</th>
<th>Observation periods</th>
<th>Class K</th>
<th>The value of $\gamma$</th>
<th>Note</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>1985-1987</td>
<td>8÷12</td>
<td>0.42±0.06</td>
<td>The period of the beginning of filling the reservoir.</td>
</tr>
<tr>
<td>2</td>
<td>1987-2014</td>
<td>8÷12</td>
<td>0.44±0.07</td>
<td>The period of the regime operation of the reservoir.</td>
</tr>
<tr>
<td>3</td>
<td>2014-2016.</td>
<td>7÷12</td>
<td>0.57±0.03</td>
<td>Study period</td>
</tr>
</tbody>
</table>

Comparison of the values of the parameters of the recurrence graph for individual intervals in the periods 1966-1973 (before construction in the area of the reservoir), 1973-1983. (during the period of construction and the beginning of filling the reservoir) and 1987-2010 (during operation) shows that the latter is characterised by a decrease in the parameter from $\gamma = 0.57$ to $\gamma = 0.34$.

Calculations were carried out using formulas (1), (2), and (3). The values of the $\gamma$ parameter obtained for different periods of averaging show that the seismic regime during the period, including the first cycle - the beginning of reservoir construction (1972-1983), differs from the second cycle - the beginning of filling the reservoir (1985-1987) and differs from the third cycle - operation of the reservoir (1987-2016).

The Packham dam of the Arsk reservoir is located on the Guzar River in an area with high seismicity - 8 points. The distance between the Gissarak and Pachkamar reservoirs is about 70 km (southwest of Gissarak). Both structures are classified as structures of the first strength class [7, 9].

Engineering seismometric observations at the facilities were organised at the end of 1986-87. (according to I.Kh. Timber and N.M. Dzhuraev [9]) and at the beginning of construction.

Since 1990, engineering seismometric observations have been improved and switched to a permanent registration scheme. The same situation was in the area of the Tupalang reservoir, with a difference of only three years before the start of the operation.

From 1990 to 2006, a new station in the area of the Pachamar reservoir was included in the observation system. Full synchronous operation of engineering seismometric observations at the facilities was organised at the end of 1986-1987 (according to I.Kh. Timbekov and N.M. Dzhuraev [9]) and at the beginning of construction.

Since 1990, engineering seismometric observations have been improved and switched to a permanent registration scheme. The same situation was in the area of the Tupalang reservoir, with a difference of only three years before the start of the operation.

From 1990 to 2006, a new station in the area of the Pachamar reservoir was included in the observation system. Full synchronous operation of all stations has been ensured since 2014.

Therefore, Table 3 shows the time intervals by dividing annual cycles according to the maximum possible representativeness. They also turned out to be more refined and generalised than in [1, 12, 23].

Another digitals / s "Pancamer Dam" has been installed in the area of the Pachkamar reservoir in the area of the dam crest. During the observation period in the Gissarak, Pachkamar, and Tupalang reservoirs, more than 1300 synchronous records of earthquakes at all stations were selected, systematised, and processed from the recorded earthquakes.

All stations have been provided since 2014.

Therefore, Table 3 shows the time intervals by dividing annual cycles according to the maximum possible representativeness. They also became more refined and generalised than in [1, 12, 23].
Another digital/s "Pancamer Dam" has been installed in the area of the Pachkamar reservoir in the area of the dam crest. During the observation period in the Gissarak, Pachkamar, and Tupalang reservoirs, more than 1300 synchronous records of earthquakes at all stations were selected, systematised, and processed from the recorded earthquakes.

Operational regime from 2010 to mid-2016 and the time of an earthquake with different energy classes in the near zone of the reservoir (Fig. 1) showed that each peak of the water pressure column corresponds to an earthquake of a certain energy class.

![Graph showing average operating regime of the Gissarak reservoir.](image)

**Fig. 1.** Average operating regime of the Gissarak reservoir (change in water height h in meters) and the time of an earthquake with different energy classes ($K = lgE$) in the near zone of the reservoir from 2010 to mid-2016 [23].

By analysing the possibility of triggered shear along the tectonic faults that are highlighted in the Gissarak Reservoir Scheme, an analytical relationship with load-discharge variation is determined.

At the end of 2015 and in the first months of 2016, there was a sharp increase in the intense water volume drop in the Gissarak reservoir. A tectonic earthquake with magnitude M=4.9 and energy class K=12.5 occurred at a depth of 10 km on 15 April 2016, 12 km west of its centre and 14 km east of the town of Kitab (coordinates =67.24; =39.06). The intensity of the shaking at the epicentre was I=6-7 magnitude (MSK). In December 2015, the Gissarak site reported the possibility of an earthquake with an energy class higher than K≥11 in the near zone of the reservoir. We classified this earthquake as induced [22,24].

This assumption is in satisfactory agreement with the results obtained in the works of D.V. Ramana and etc. [13], J.M. Ferreira and etc. [17], G.Kaufmann, F.Amelung [18], A.Kangi, N.Heidari [20], H.Ping, H.Yuliang [25] in Koina and Varna reservoirs (India), Asu (Brazil), Mid (USA), Karun III (Iran) and in zones of big water reservoirs of China under completely different seismotectonic conditions.

### 4 Conclusions

Analysis of the total and background seismicity of the Hissarak, Pachkam Ar and Tupalang reservoir zones in Southern Uzbekistan has shown that the distribution of local earthquake focuses in the study area is different for different years. According to materials from 2007 to 2014 with energy class K≥8.0 and in recent years from 2014 to June 2016 with K≥7.0,
there is a strong variation in the number of earthquakes and densification of epicenters in narrow zones that coincide with the strike direction of local fault structures in the seismic regime.

A unique catalogue of local earthquakes in their influence zone has been compiled by tracking the seismicity levels of the Gissarak, Pachkam Ar and Tupalang reservoirs in southern Uzbekistan. An important feature of the created catalogue of local earthquakes is its comparability both with regional catalogues and transboundary states' catalogues.

Tracking of local seismicity levels of reservoir zones in South Uzbekistan has revealed several evidence of the negative impact of large hydro-technical reservoirs on deformation and seismic state of zones of the active impact of reservoirs.

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